

# Climate Change Fingerprinting

with help from chatGPT

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## 1 Classical Climate Change Fingerprinting

### 1.1 Set up the linear model

$$\mathbf{y} = \mathbf{X}\boldsymbol{\beta} + \boldsymbol{\varepsilon}, \tag{1}$$

- $\mathbf{y}$  —  $N$ -vector of observed anomalies (e.g. gridded temperatures concatenated over space  $\times$  time).
- $\mathbf{X}$  —  $N \times J$  design matrix whose  $J$  columns are the *fingerprints* (model-simulated responses) of the external forcings of interest: greenhouse gases, aerosols, solar, volcano, . . .
- $\boldsymbol{\beta}$  — unknown *scaling factors* (one per forcing) to estimate and test.
- $\boldsymbol{\varepsilon}$  — internal variability, assumed mean 0 with covariance  $\boldsymbol{\Sigma} = \text{Cov}[\boldsymbol{\varepsilon}]$ .

Because  $\mathbf{y}$  has strong space–time correlations, ordinary least squares (OLS) would underestimate uncertainty. Generalized least squares (GLS) fixes this by explicitly using  $\Sigma$ .

## 1.2 Estimate the noise covariance matrix

Take  $\Sigma$  from long control-run simulations, filtered exactly like the observations:

$$\hat{\Sigma} = \frac{1}{M} \sum_{m=1}^M (\mathbf{c}^{(m)} - \bar{\mathbf{c}})(\mathbf{c}^{(m)} - \bar{\mathbf{c}})^\top. \quad (2)$$

Large  $N$  makes  $\hat{\Sigma}$  ill-conditioned, so in practice we

- rotate into the leading EOFs of  $\hat{\Sigma}$  and truncate, and/or
- apply shrinkage or ridge regularisation.

## 1.3 Pre-whiten

Compute the symmetric square-root inverse  $\hat{\Sigma}^{-1/2}$  (in the retained EOF subspace) and premultiply:

$$\tilde{\mathbf{y}} = \hat{\Sigma}^{-1/2} \mathbf{y}, \quad \tilde{\mathbf{X}} = \hat{\Sigma}^{-1/2} \mathbf{X}.$$

After whitening the noise is (approximately) i.i.d. and we may use OLS on  $(\tilde{\mathbf{y}}, \tilde{\mathbf{X}})$ . In compact form

$$\boxed{\hat{\beta}_{\text{GLS}} = (\mathbf{X}^\top \hat{\Sigma}^{-1} \mathbf{X})^{-1} \mathbf{X}^\top \hat{\Sigma}^{-1} \mathbf{y}} \quad (3)$$

with  $\text{Var}(\hat{\beta}) = (\mathbf{X}^\top \hat{\Sigma}^{-1} \mathbf{X})^{-1}$ .

## 1.4 Detection and attribution

- 1) **Detection:** For each forcing  $j$ , test  $H_0: \beta_j = 0$ . Rejecting (e.g. at 5%) means the fingerprint is *detected*.
- 2) **Attribution / consistency:** Check whether  $\beta_j$  is statistically consistent with 1. If 1 lies in the confidence interval, the model response amplitude is consistent with observations.

### 3) Residual consistency:

$$Q = \frac{\mathbf{r}^\top \hat{\Sigma}^{-1} \mathbf{r}}{N - J}, \quad \mathbf{r} = \mathbf{y} - \mathbf{X}\hat{\boldsymbol{\beta}}.$$

Compare  $Q$  with the corresponding  $\chi^2$  quantile; failure suggests a missing forcing, non-linear response, or mis-specified  $\hat{\Sigma}$ .

## 1.5 Practical refinements

Issue	Typical remedy
Errors in fingerprints $\mathbf{X}$ (finite ensemble size)	Total-least-squares (TLS), <i>generalised</i> TLS, or recent estimating-equation approaches.
Multicollinearity between forcings	Dimensionality reduction (eigen-fingerprints), ridge regression, Bayesian priors.
High-dimensional $\Sigma$	Shrinkage estimators, covariance tapering, or projection onto a reduced EOF basis before GLS.
Low signal-to-noise	Optimal filtering: weight fingerprints toward low-noise directions (the original “optimal fingerprint” idea).

## 1.6 Readings

- Hegerl, G. C. *et al.* (1996), *Journal of Climate*.
- Hannart, A. (2019), *ASCMO* — explanation of projection choices and S/N optimization.

## 1.7 Key take-aways

- Classical fingerprinting is *just* a GLS regression whose design matrix is built from model-simulated responses, and whose noise covariance comes from control runs.

- The estimated *scaling factors* quantify how strongly each forcing contributes to the observed change; their confidence intervals provide the detection/attribution statements.
- Everything hinges on a realistic  $\Sigma$ ; dimension reduction and regularisation are therefore essential in real applications.

## 1.8 Covariance Matrix Estimation

Accurate estimation of the internal variability covariance  $\Sigma$  is crucial especially when it needs to be inverted. First, the Sample Covariance Matrix is calculated as

$$C = \frac{1}{N-1} \sum_{i=1}^N (\mathbf{z}_i - \bar{\mathbf{z}})(\mathbf{z}_i - \bar{\mathbf{z}})^T.$$

This matrix is often unstable or singular when the number of variables (grid points)  $p$  is large relative to the sample size (number of months)  $N$ . To address this, we can use a Shrinkage Estimator,

$$\widehat{\Sigma} = \lambda \mathbf{T} + (1 - \lambda)C, \quad \lambda \in [0, 1],$$

where

- $\mathbf{T}$ : stable target (e.g. scaled identity or diagonal).
- $\lambda$ : shrinkage intensity.

This guarantees positive-definiteness, reduces estimation error in high dimensions. One specific variant is the Ledoit–Wolf Optimal Shrinkage, where,

$$\lambda^* = \frac{\mathbb{E}[\|\mathbf{S} - \Sigma\|_F^2]}{\mathbb{E}[\|\mathbf{T} - \Sigma\|_F^2 + \|\mathbf{S} - \mathbf{T}\|_F^2]},$$

estimated with plug-in values from the data.

Other Estimators are Tapering/Banding which impose distance-based decay of covariances, Factor Models, which use a low-rank signal plus white noise, or Sparse Precision Estimators (graphical lasso) which estimate a sparse inverse covariance (conditional dependencies).